

Probabilistic Networks models for Statistical Seismology

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Questions

- Can we identify local spatial correlation structure in seismic catalog data?
- Given an earthquake, which are the areas where the aftershocks could be worse?
- Does knowing the seismic activity in a given location help to explain the seismic activity in nearby locations?
- Does seismic activity exhibit long-distance connections between regions that cannot be explained only by local behavior?

Modeling relationships in Seismology (and other fields)

- Correlation: basic measure of linear pairwise relationships
- Covariance matrix Σ : collection of relationships
- Estimates of Σ required in procedures such as PCA, CCA, MANOVA, etc.
- Estimating (functions of) Σ and $\Omega = \Sigma^{-1}$ are of statistical interest
- Estimating Σ is difficult in high dimensions

- Matrix Σ or Ω of size p -by- p has $O(p^2)$ elements
- Estimating $O(p^2)$ parameters with classical estimators is not viable, especially when $n \ll p$
- **Reliably estimate small number of parameters in Σ**
- Model selection: zero/non-zero structure recovery
- Gives rise to sparse estimates of Σ or Ω
- Sparsity pattern can be represented by graphs/networks

- **Many physical networks are assumed to be sparse**
- Complex networks (internet, citation networks, social networks) tend to be sparse [Newman, 2003]
- Genetic networks are sparse [Gardner et al, 2003, Jeong et al, 2001]
- Model selection: recovery of edges in graphs

Gaussian Graphical Models (GGM)

- Assume $Y = (Y_1, \dots, Y_p)'$ has distribution $N_p(0, \Sigma)$
- Denote $V = \{1, 2, \dots, p\}$
- Covariance matrix $\text{cov}(Y) = \Sigma$ encodes marginal dependencies

$$Y_i \perp Y_j \iff \text{cov}(Y_i, Y_j) = [\Sigma]_{ij} = 0$$

- Inverse covariance matrix $\Omega = \Sigma^{-1}$ encodes conditional dependencies given the rest

$$\underbrace{(Y_i \perp Y_j \mid Y_{V \setminus \{i,j\}})}_{\text{conditional independence}} \iff \underbrace{[\Omega]_{ij} = 0}_{\text{matrix element}}$$

- Also known as Markov Random Fields (MRF)

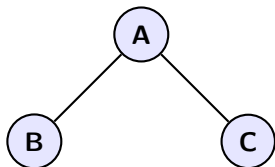
Gaussian Graphical Models (GGM)

- Graph summarizes relationships with nodes $V = \{1, \dots, p\}$ and set E of edges

$$\underbrace{[\Omega]_{ij} = 0}_{\text{matrix element}} \iff \underbrace{i \not\sim j}_{\text{network/graph}}$$

- Build a graph from sparse Ω

$$\Omega = \begin{pmatrix} & \text{A} & \text{B} & \text{C} \\ \text{A} & 1 & 0.2 & 0.3 \\ \text{B} & 0.2 & 2 & 0 \\ \text{C} & 0.3 & 0 & 1.2 \end{pmatrix}$$

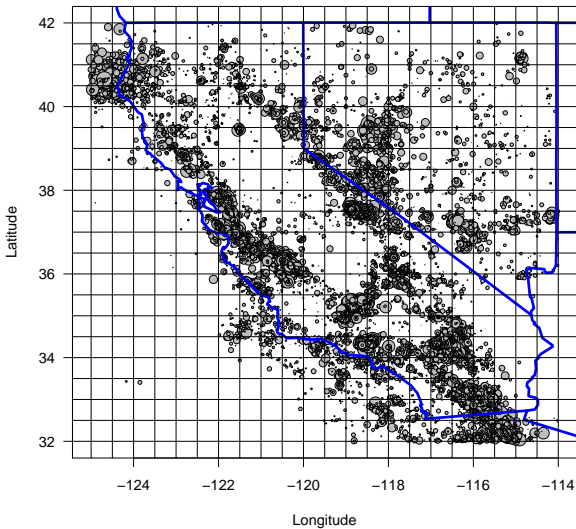


Description of the Data

We extracted data from the Advanced National Seismic System (ANSS) catalog.

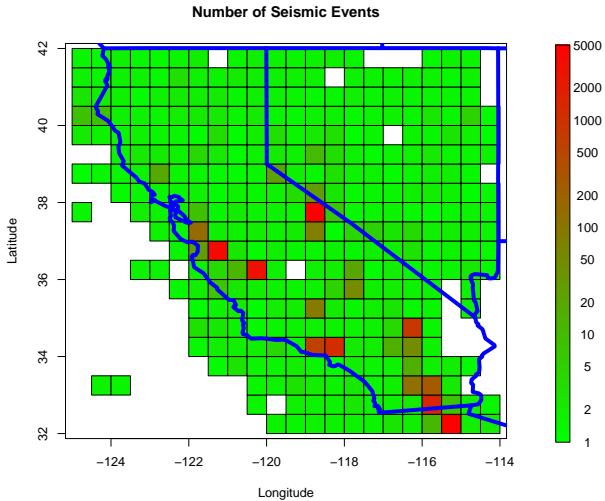
- The data contains all seismic events of magnitude at least 3.0 in an $11^\circ \times 10^\circ$ rectangle covering California and Nevada since January 1, 1950.
- We grouped the data into $0.5^\circ \times 0.5^\circ$ grid boxes and computed (1) the cumulative magnitudes and (2) the number of all events in a six month period in each grid box.
- Regarding the gridded cumulative magnitudes (or the number of events) during a given six month period as an observation from a random vector, we have a random vector of length $p = 340$ and a total of $n = 127$ observations.

Plot of the Data



Larger points denote seismic events of higher magnitude.

Seismic Events in Each Geographic Bin



Methodology

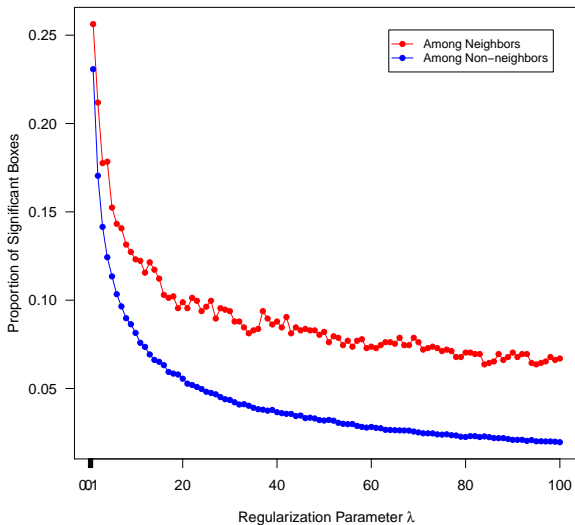
- Neighboring grid boxes were defined to be those boxes directly adjacent to a given grid box.
- We computed the sample correlation matrix between grid boxes.
- Since $p > n$, the sample correlation matrix is a poor estimate of the true correlation matrix.
- Also, since $p > n$, the sample correlation matrix is singular. This means that its inverse (the sample partial correlation matrix) cannot be computed at all.
- Some form of regularization is required to obtain good estimates.

Regularization Approaches Considered

- **Correlation hard-thresholding** estimates the correlation matrix. We can consider either the Pearson correlation or the Spearman rank correlation.
- **Graphical LASSO (GLASSO)** estimates the partial correlation matrix. Again, we can consider either the Pearson correlation or the Spearman rank correlation.
- **SPACE** also estimates the partial correlation matrix. Here we can only consider the Pearson correlation.

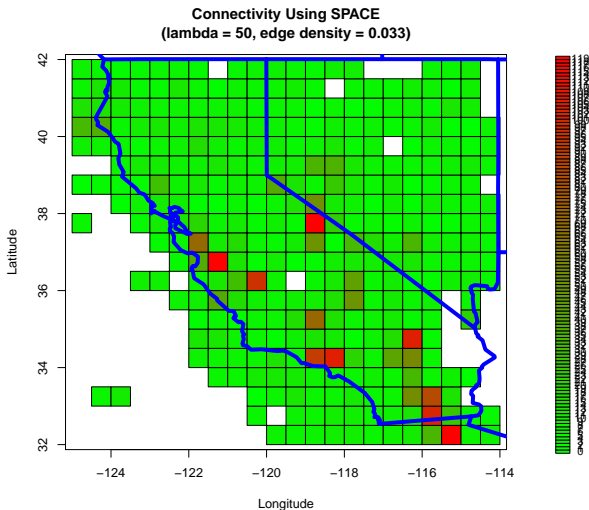
SPACE: Sum of Magnitudes of Seismic Events

Assuming 8 Adjacent Boxes



- Like the GLASSO, SPACE also estimates the partial correlation matrix, the inverse of the correlation matrix itself.
- Once again, we still see the same phenomenon: nodes that are geographic neighbors are much more likely than non-neighbors to have a nonzero element (here, a nonzero partial correlation).
- Again, the other plots also behave similarly (next few slides).

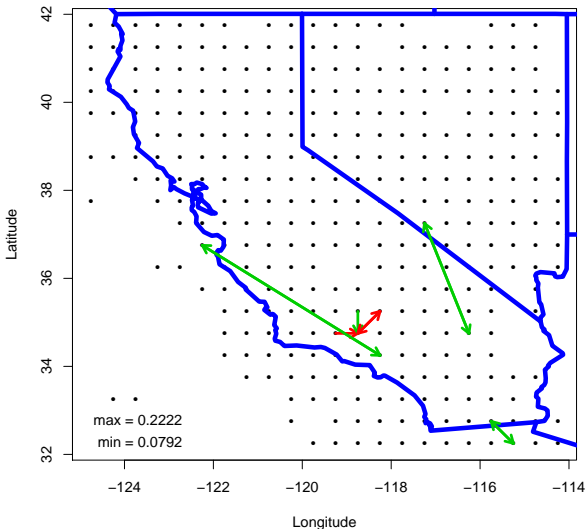
SPACE: Sum of Magnitudes of Seismic Events



The **connectivity** of node i is the number of other nodes for which the estimated correlation with node i is nonzero.

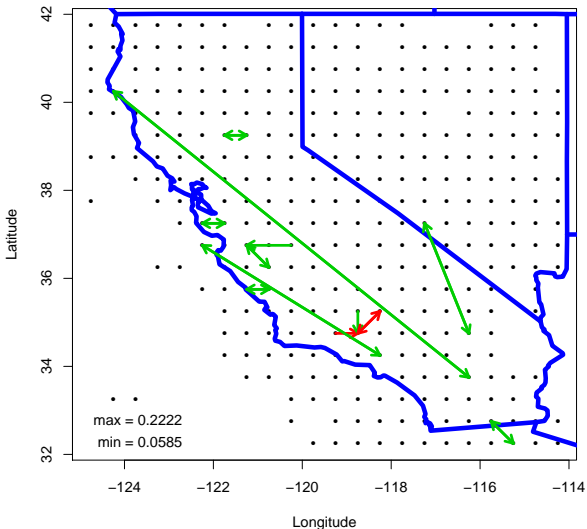
SPACE: Sum of Magnitudes of Seismic Events

Strongest Partial Correlations for Each Node (Top 10 and ties;
SPACE with $\lambda = 150$, edge density = 0.016)



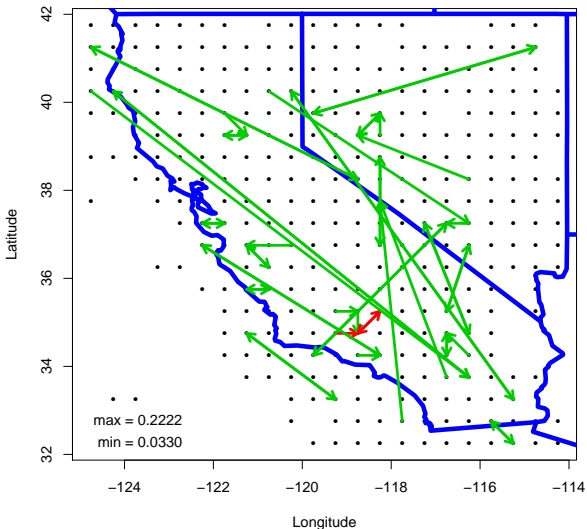
SPACE: Sum of Magnitudes of Seismic Events

Strongest Partial Correlations for Each Node (Top 20 and ties;
SPACE with $\lambda = 150$, edge density = 0.016)

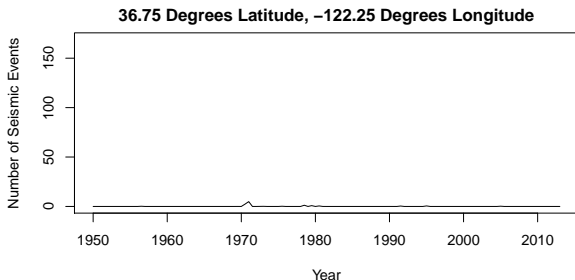
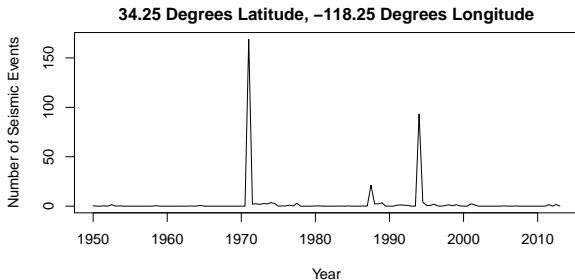


SPACE: Sum of Magnitudes of Seismic Events

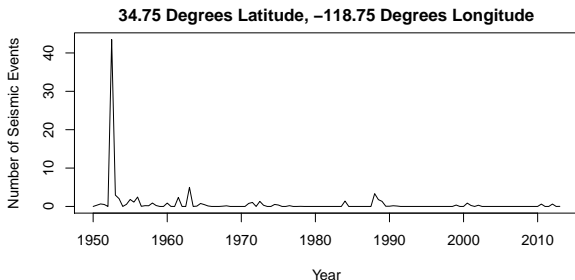
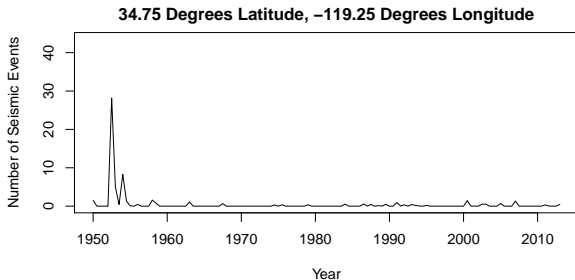
Strongest Partial Correlations for Each Node (Top 50 and ties;
SPACE with $\lambda = 150$, edge density = 0.016)



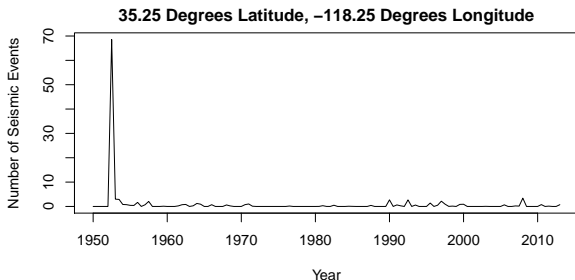
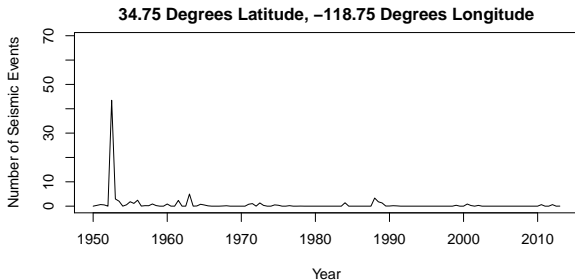
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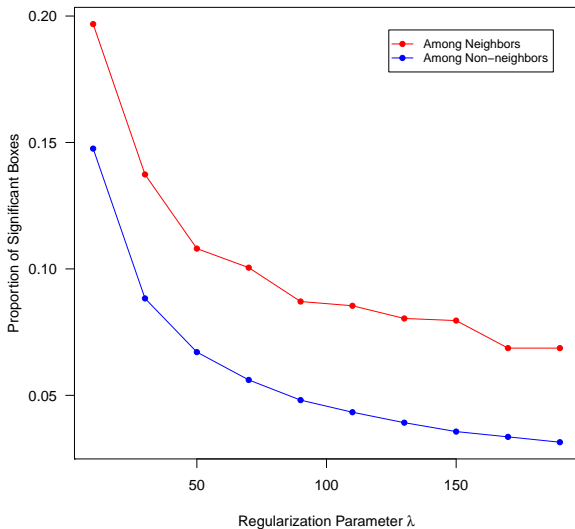


SPACE: Number of Seismic Events

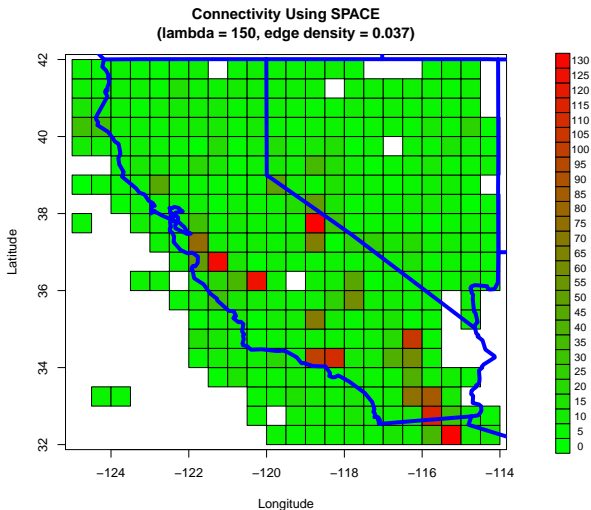
- The results are fairly similar if we consider simply the number of seismic events in each spatial and temporal bin, rather than the sum of the magnitudes.
- This alternative approach yields the plots shown in the following slides.

SPACE: Number of Seismic Events

Assuming 8 Adjacent Boxes



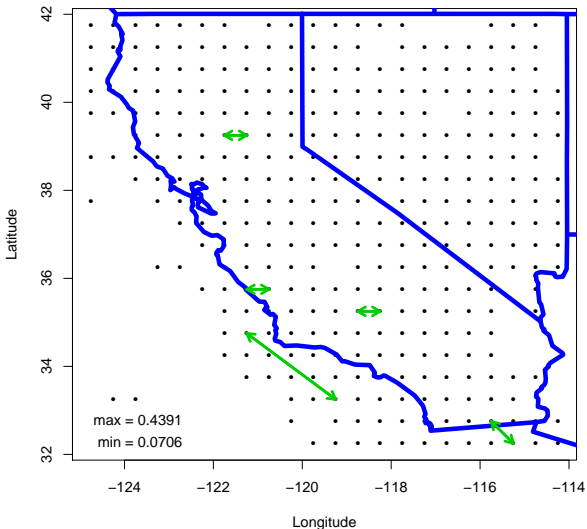
SPACE: Number of Seismic Events



The **connectivity** of node i is the number of other nodes for which the estimated correlation with node i is nonzero.

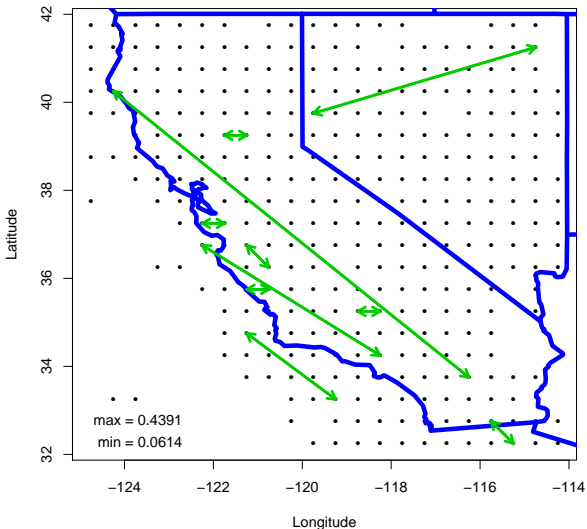
SPACE: Number of Seismic Events

Strongest Partial Correlations for Each Node (Top 10 and ties;
SPACE with $\lambda = 150$, edge density = 0.037)



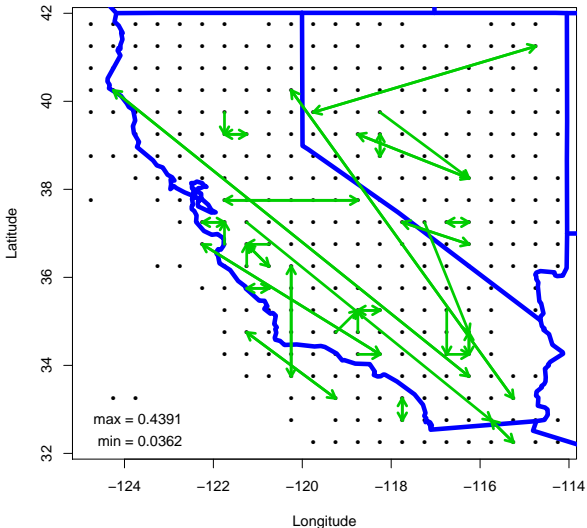
SPACE: Number of Seismic Events

Strongest Partial Correlations for Each Node (Top 20 and ties;
SPACE with $\lambda = 150$, edge density = 0.037)

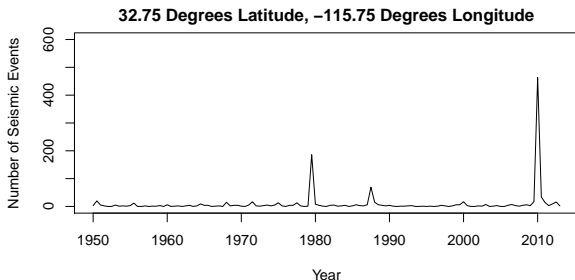
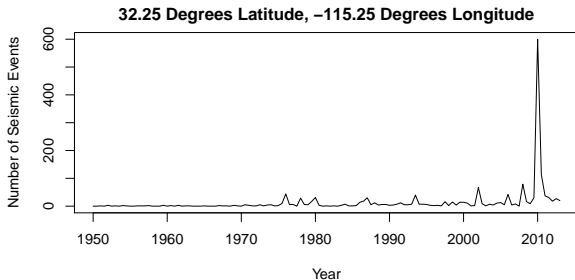


SPACE: Number of Seismic Events

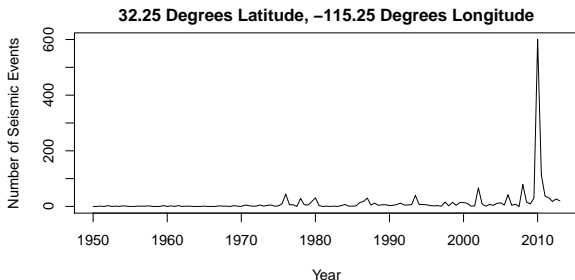
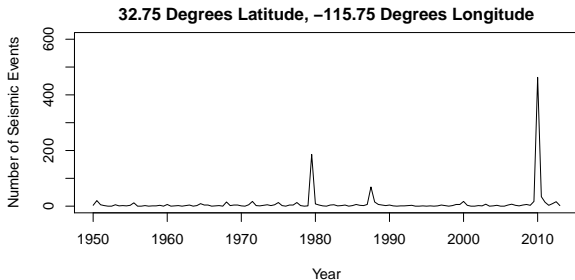
Strongest Partial Correlations for Each Node (Top 50 and ties;
SPACE with $\lambda = 150$, edge density = 0.037)



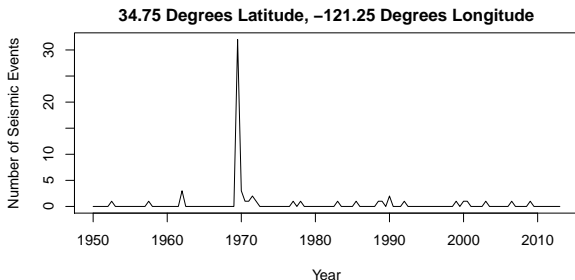
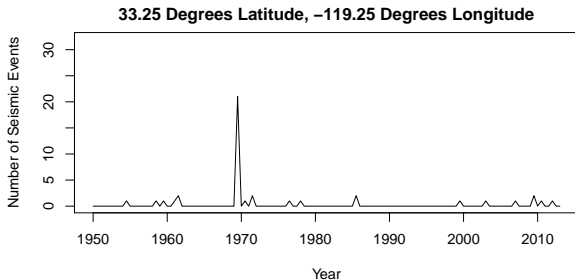
SPACE: Number of Seismic Events



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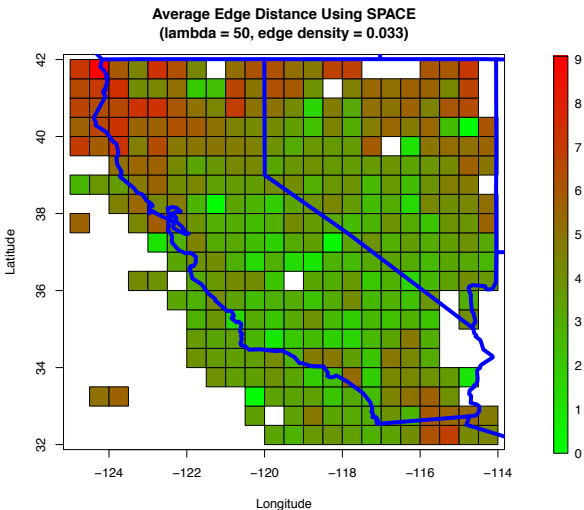


SPACE: Number of Seismic Events



Conclusions

- Data suggests that Markov random fields can be useful for describing seismic activity.
- Seismic activity in a target area is more related to activity in geographically neighboring areas than to activity in non-neighboring areas.
- However, long-distance connections can still have substantial explanatory value when they are present.



The **average edge distance** for node i is its average distance from the other nodes with which it has a nonzero estimated correlation.