Blending Ensembles of Regional Climate Model Predictions

Bruno Sansó

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The Economist

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It's hard to make predictions - especially about the future.

Yogi Berra

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Regional Climate Models (RCM) produce a "dynamic downscaling" of the output of GCMs. They simulate relatively short-term atmospheric and land-surface processes and the interactions between the two, at a spatial resolution of about 50 km.

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In this talk we focus on the multi-model uncertainty.

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The traditional setting for assessment and calibration of a computer model assumes that both, model and observations provide information about a true, unobserved quantity, say ξ . Then

$$Y = \xi + \varepsilon$$
, and $F(\boldsymbol{\theta}) = \xi + \delta$

where ε is observational error and δ is model discrepancy.

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- Estimation of θ provides a calibration of the model.
- Estimation of ξ provides information about the property of interest, using both simulations and observations.

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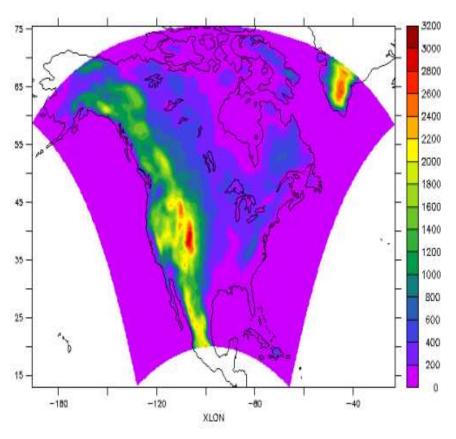
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To tackle this issue we can:

- Average over large areas and time spans.
- Consider large scale summaries of the spatial and temporal fields, i.e. trends, cycles, patterns, indexes.
- Use Space-time models for smoothing.

REGIONAL MODELS

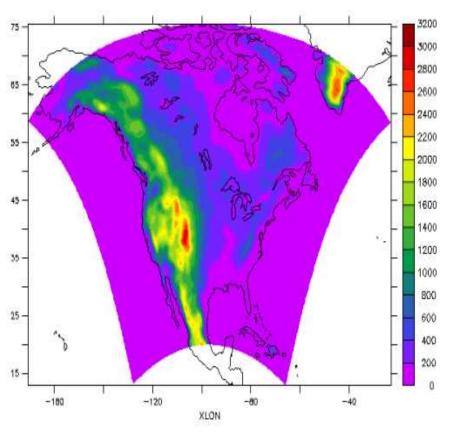
NARCCAP DOMAIN



• NARCCAP is a program to produce high resolution climate change simulations over the US, Canada and Mexico.

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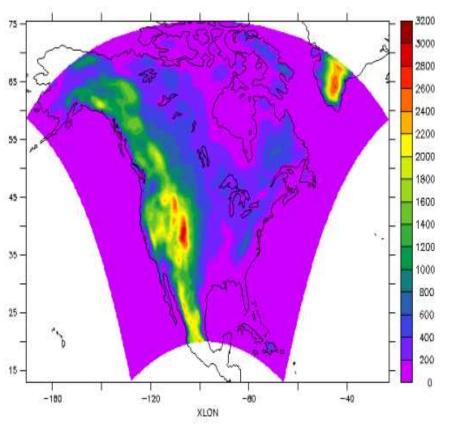
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- The goal is to assess climate variability at a regional level.
- All RCMs use the same 50 km resolution and the same future emission scenario (A2)

NARCCAP COMBINATIONS

• NARCCAP considers six different RCMs, four different AOGCM, NCEP reanalysis and two time slices.

NARCCAP Combinations

AOGCMs

RCMs	GFDL	HADCM3	CGCM3	CCSM
RegCM3	X		X	
ECPC	X	X		
PRECIS	X	X		
CRCM			X	X
WRF			\mathbf{x}	X
MM5		X		X

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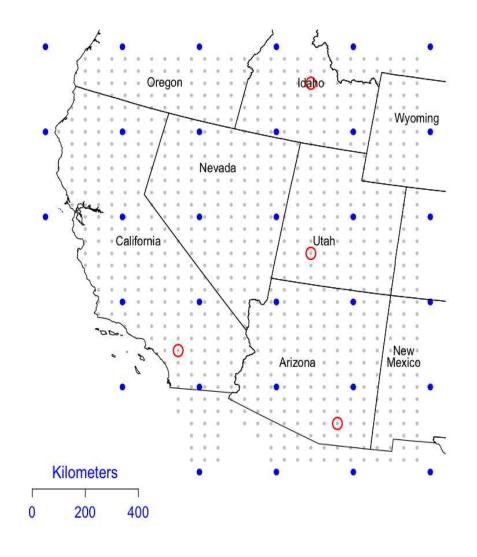
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- Not every combination of the RCMs and the AOGCM are considered, so the experiment resulted in a fractional factorial design.
- All models consider present day conditions from 1971 to 2000 and future simulations 2041 to 2070.

NARCCAP Combinations

AOGCMs

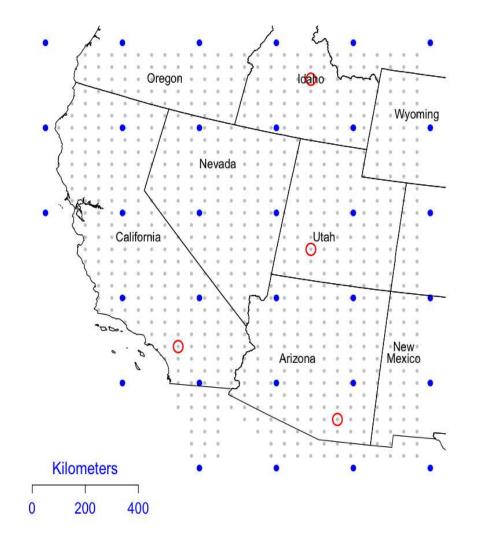
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The Domain of our Analysis



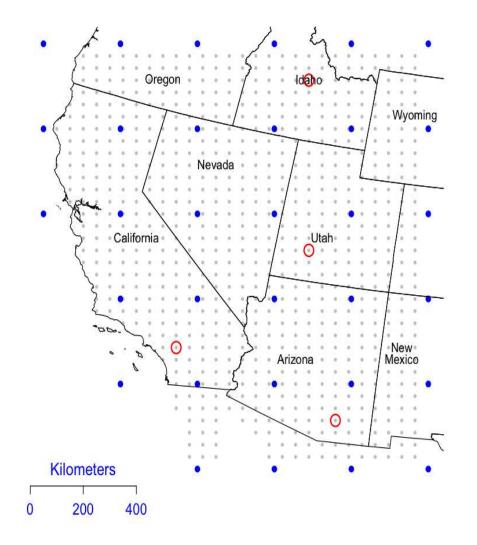
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The Domain of our Analysis



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- We consider the simulations obtained using RegCM3 under NCEP, GFDL and CGCM3 forcings.
- We study the variability of yearly mean summer temperature at each of the 802 locations.

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- To obtain interpolated fields we processed 3 hour data by detrending using location and elevation. We then estimated exponential variogram parameters with nugget.
- We did a simple kriging of the residuals and then averaged of all the 3 hourly values.

GOALS

• Validate the RCM simulations with respect to the observational records.

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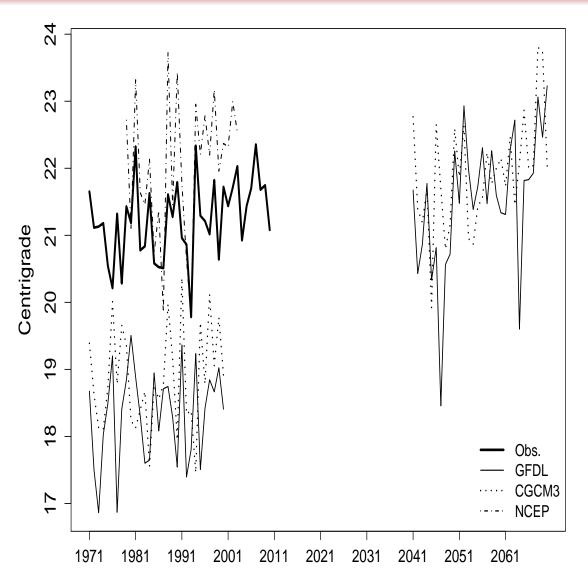
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- Explore trends of spatial and temporal variability that are common for the four sources of information.

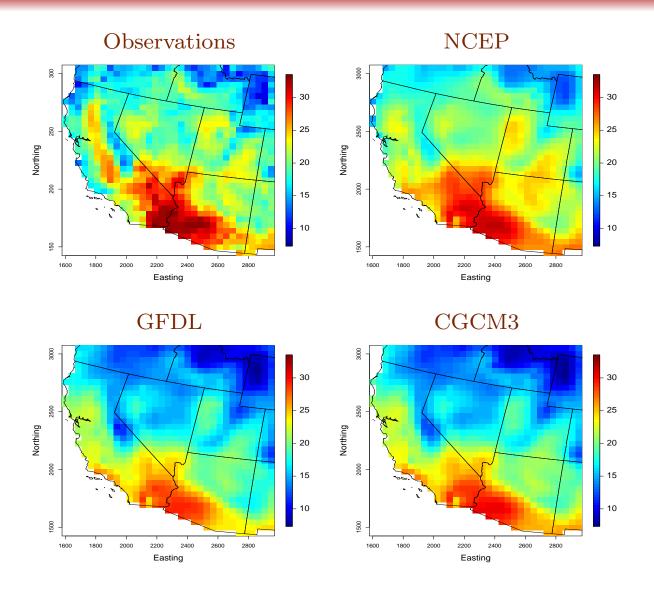
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- Merge the four data sources to obtain blended reconstructions and forecasts, including probabilistic measures of uncertainty.

Averages Over Space



Note the gap in the first part of the 21st Cent., and the discrepancy between obs. and simul. during the 20th Cent. for CGCM3 and GFDL.

Averages Over Time



Our Model

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All four data sources correspond to a common space-time process. RCMs deviations from that process are time and space varying.

$$y_{t}(s) = x_{t}^{T}(s)\eta + \xi(t - t_{0}) + \omega_{t}(s) + + \epsilon_{t}(s)$$

$$y_{jt}^{CM}(s) = x_{t}^{T}(s)\eta + \xi(t - t_{0}) + \omega_{t}(s) + d_{jt}(s) + \epsilon_{jt}(s)$$
covariates trend baseline discrepancy

 $\epsilon_t(s)$ and $\epsilon_{jt}(s)$ are observational errors.

The dimensionality of $\omega_t(s)$ is reduced with a predictive Gaussian process approach:

$$\omega_t(\mathbf{s}) = \sum_{m=1}^{M} B_m(\mathbf{s}) \gamma_{m,t} + \tilde{\varepsilon}_t(\mathbf{s}) = \mathbf{B}(\mathbf{s})^T \boldsymbol{\gamma}_t + \tilde{\varepsilon}_t(\mathbf{s})$$

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$$B_m(\mathbf{s}) = [\mathbf{v}(\mathbf{s})^T \mathbf{H}^{-1}]_m, \ \boldsymbol{\gamma}_t \sim N(\varphi \boldsymbol{\gamma}_{t-1}, \mathbf{H}) \text{ and }$$

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$$v(s) = \tau^2(\rho(s, s_1^*; \phi), \dots, \rho(s, s_M^*; \phi)) \text{ and } H_{lk} = \tau^2 \rho(s_l^*, s_k^*; \phi).$$

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regular grid with a resolution of 290 km.

TIME EVOLUTION

Consider the spectral decomposition $\mathbf{H} = \mathbf{P} \mathbf{\Lambda} \mathbf{P}^T$, \mathbf{P} orthogonal and $\mathbf{\Lambda}$ diagonal. Let $\gamma_t = \mathbf{P} \alpha_t$, $\forall t$, then

$$\omega_t(s) = \boldsymbol{B}(s)^T \boldsymbol{P} \boldsymbol{\alpha}_t = \boldsymbol{\psi}(s)^T \boldsymbol{\alpha}_t \text{ and } \boldsymbol{\alpha}_t \sim N(\varphi \boldsymbol{\alpha}_{t-1}, \boldsymbol{\Lambda}).$$

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A similar representation for d_{jt} yields

$$\omega_t(s) + d_{jt}(s) = \boldsymbol{B}(s)^T (\boldsymbol{\gamma}_t + \boldsymbol{\gamma}_{jt}) = \boldsymbol{\psi}(s)^T (\boldsymbol{\alpha}_t + \boldsymbol{\alpha}_{jt}),$$

and

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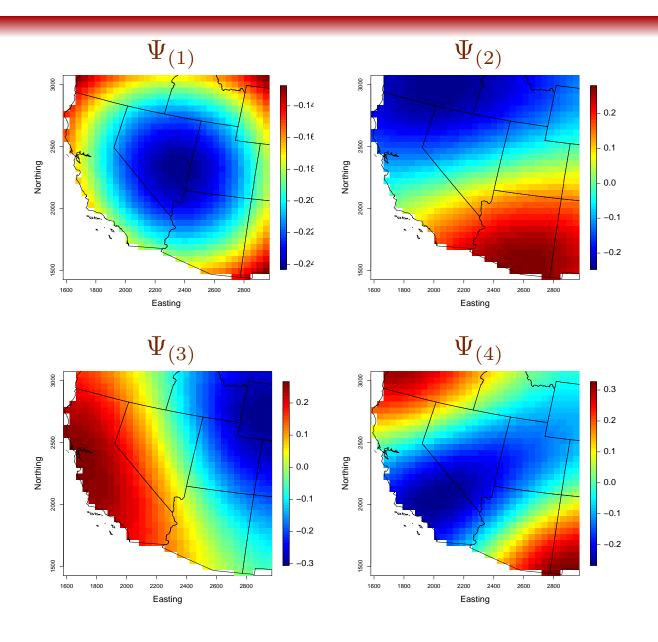
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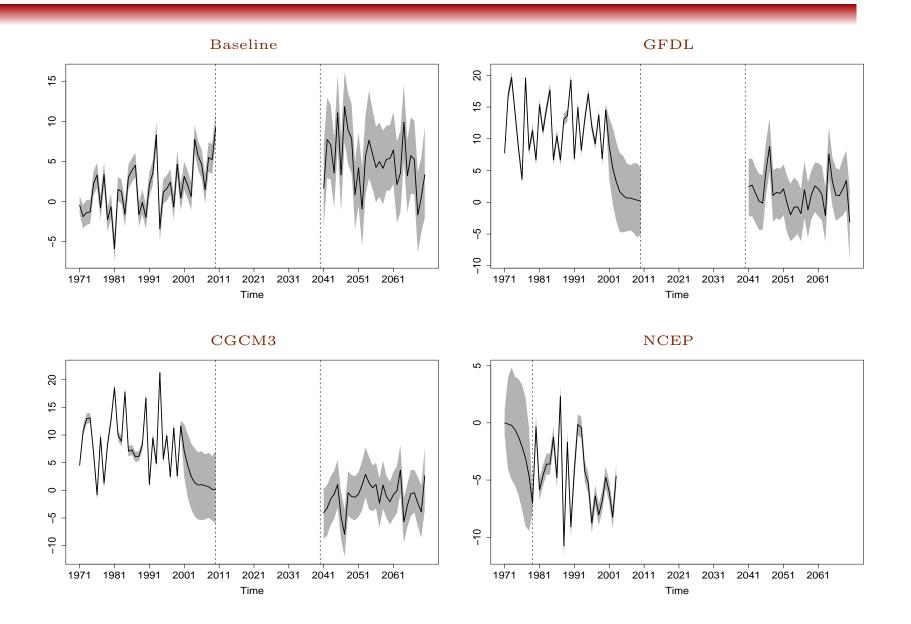
$$\boldsymbol{\alpha}_{jt} \sim N(\varphi_j \boldsymbol{\alpha}_{j,t-1}, \boldsymbol{\Lambda}_j).$$

The fields $\psi_m(s)$ are not orthogonal, but the corresponding coefficients are independent with decreasing variance.

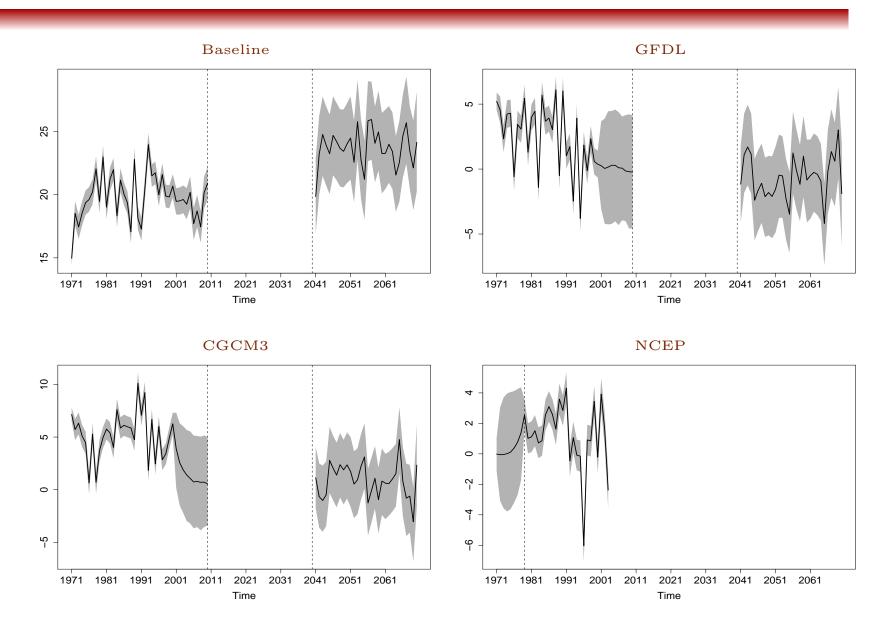
FIRST FOUR FACTORS



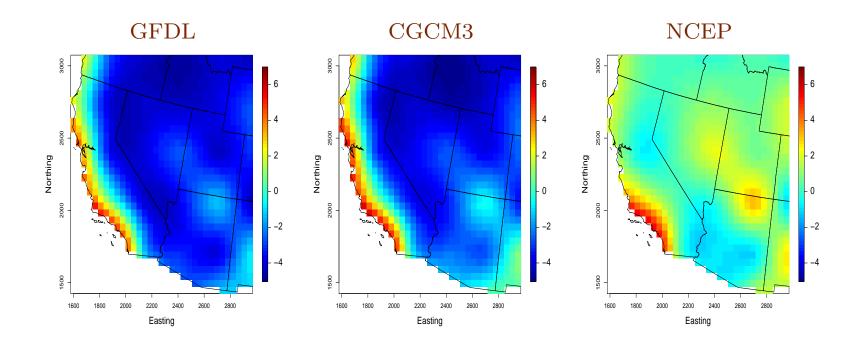
FIRST FACTOR COEFFICIENTS



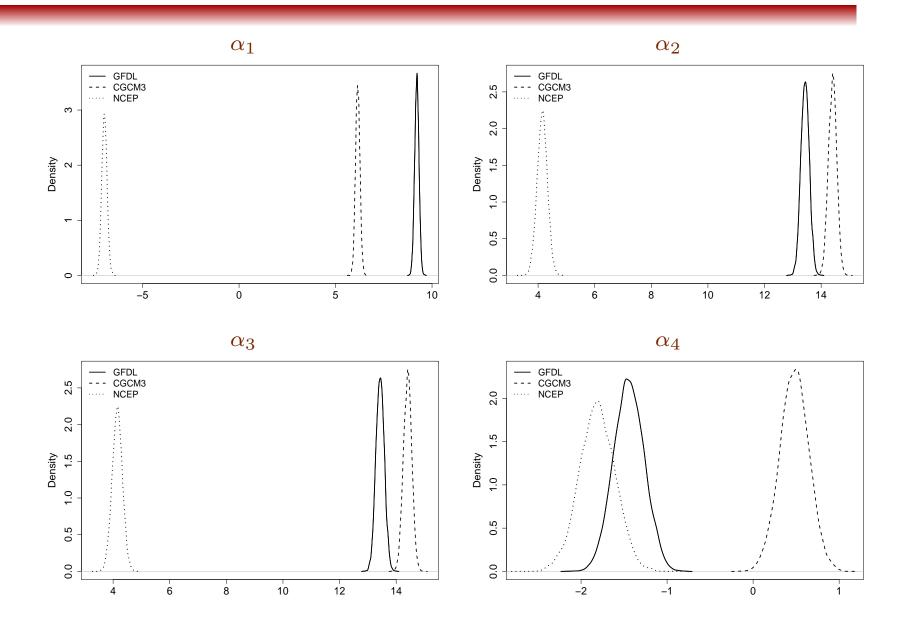
SECOND FACTOR COEFFICIENTS



Constant Discrepancy Model



CONSTANT DISCREPANCY COEFFICIENTS



Model Assessment

We assess our future predictions by taking a training set (1971–1990) and a test set (1991–2000). We consider:

- Continuous rank probability scores.
- Energy scores
- Root mean square error
- Mean absolute error

Model Assessment

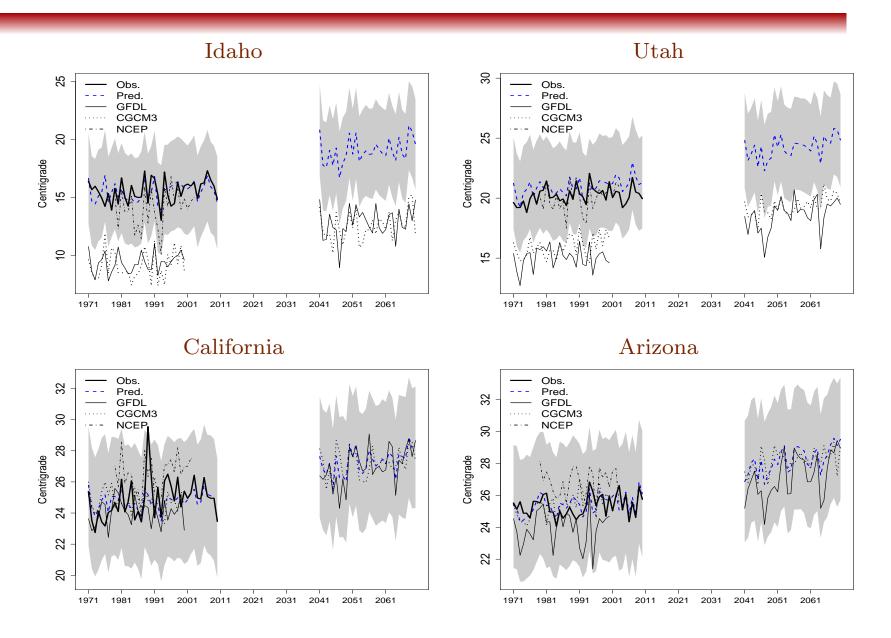
Forecast	CRPS	ES	RMSE	MAE
CGCM3	2.91	108.80	3.98	3.38
GFDL	3.20	115.50	4.17	3.67
Merged	3.01	110.40	4.06	3.51
Observations	0.59	20.15	1.03	0.79
Model 1 (32 knots)	2.19	74.20	3.53	2.95
Model 2 (32 knots)	1.20	42.94	2.14	1.69
Model 2 (68 knots)	1.22	43.72	2.17	1.70

Model Assessment

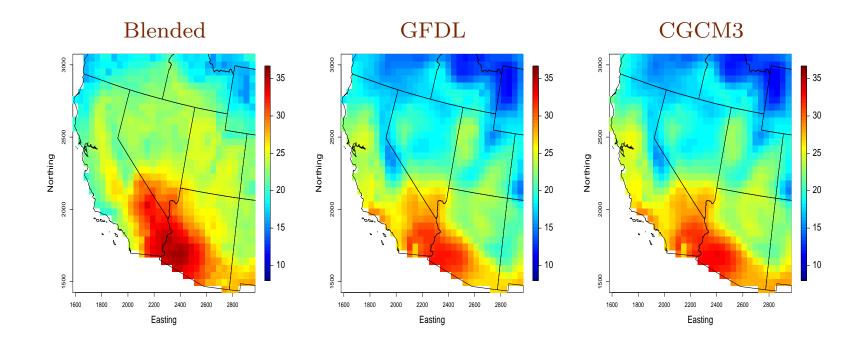
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All three statistical procedures improve the predictions of the model runs. The best method is obtained with a coarse grid and constant discrepancies.

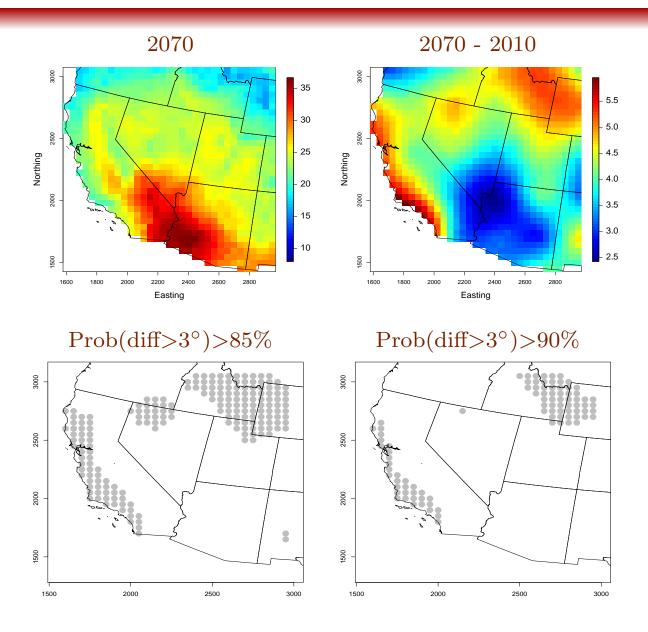
PREDICTED TIME SERIES



Average Predictions 2041–2070



BLENDED PREDICTIONS



• We use hierarchical models to compare and blend information from different climate model simulations and obtain unified predictions.

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- Our model provides a quantification of the uncertainties associated with the predictions.
- We use spatial and temporal models to introduce smoothing in time and space.
- Our spatial factor model reduces computations and allows for the description of patterns, cycles and trends that can be used as summaries of the analysis.

REFERENCES

- Esther Salazar, Bruno Sansó, Andrew Finley, Dorit Hammerling, Ingelin Steinsland, Xia Wang and Paul Delamater (2011) "Comparing and Blending Regional Climate Model Predictions for the American Southwest". To appear in the Journal of Agricultural Biological and Ecological Statistics.
- Francisco Beltrán, Bruno Sansó, Ricardo Lemos and Roy Mendelssohn (2011) "Joint Projections of North Pacific Sea Surface Temperature from Different Global Climate Models".
 Tech Report UCSC-SOE-11-03.

REFERENCES

- Claudia Tebaldi and Bruno Sansó (2009) "Joint Projections of Temperature and Precipitation Change from Multiple Climate Models: A Hierarchical Bayes Approach" *Journal of the Royal Statistical Society Series A*, vol. 172, pp. 83–106.
- Mark A. Snyder, Bruno Sansó and Lisa C. Sloan (2007).
 "Validation of Climate Model Output using Bayesian Statistical Methods". *Climatic Change*, vol 83, pp 457–476, 10.1007/s10584-007-9262-3.
- Bruno Sansó and Lelys Guenni (2004) "A Bayesian Approach to Compare Observed Rainfall Data to Deterministic Simulations", *Environmetrics*, vol 15, pp. 597–612.